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Phytotoxin sorption to clay minerals



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Abstract

Background: Phytotoxins of various classes and origin are often found in their cationic form in the soil environment and thus, their overall soil behavior may be strongly affected by all geosorbents presenting cation exchange capacity (CEC). In addition to soil organic carbon (SOC), clays may exhibit great potential as sorbents for cationic organic chemicals. Therefore, 52 compounds of the major phytotoxin classes alkaloids, terpenoids and steroids were investigated with regard to their sorption behavior to the clay minerals kaolinite (low CEC) and montmorillonite (high CEC) by means of continuous flow column sorption experiments as a high-throughput alternative to traditional batch sorption experiments.

Results: In total, sorption coefficients $\log D_{\rm clay}$ [L kg⁻¹] were quantifiable for 26 phytotoxins on kaolinite ($\log D_{\rm clay} > 0.1$) and 33 on montmorillonite ($\log D_{\rm clay} > 0.5$). They ranged from 0.14 ± 0.09 for the pyrrolizidine alkaloid senkirkine on kaolinite to 3.05 ± 0.03 for the indole alkaloid brucine on montmorillonite. Although maximum sorbed concentrations lay well below the CEC for both clay minerals, sorption non-linearity was observed in some cases where as little as 0.1% of all cation exchange sites were occupied. Contrary to the expectations, sorption non-linearity could not be wholly explained by saturation of available sorption sites; for protonated tertiary amines with aromatic moieties, cooperative sorption seemingly took place and the results indicated a significant increase in sorption affinities within a very limited concentration range. Comparing montmorillonite and SOC, notable differences in preferences of cationic sorbates were observed between phytotoxins with and without aromatic moieties (e.g., isoquinoline versus pyrrolizidine alkaloids) as well as between N-heterocycles and N-heteroaromatics in particular (e.g., strychnine versus gramine; both indole alkaloids).

Conclusions: Overall, clay sorption seems a result of the interplay of charge location on the sorbent and various structural features of the sorbates. To confirm observed tendencies towards cooperative sorption for certain cationic phytotoxins, further studies with higher concentrations are needed. Nevertheless, obtained sorption coefficients indicate that a high proportion of phytotoxin sorption in soils may be attributed to clay minerals. Thus, clay minerals possess the ability to decrease total cationic phytotoxin environmental mobility.

Keywords: Natural toxins, Aluminosilicates, Montmorillonite, Kaolinite, Freundlich isotherm, Cooperative sorption, Environmental mobility

Background

A tremendous number of structurally different secondary metabolites are produced by plants [1, 2] and subsequently emitted to environmental compartments such

as air, soil and water. Among other functions, they can seemingly serve as signal molecules for, e.g., pollinators or as defense compounds (phytotoxins) against, e.g., pathogens or insects. As such, they are important for the plants' survival [3]. Phytotoxins are biologically active and may also pose a risk to non-target organisms [4–6]. As "nature's own pesticides" [4], phytotoxins are also used as biopesticides [7, 8] or serve as model compounds for the development of such chemicals [9]. Many phytotoxins are protonated under environmentally relevant

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conditions; particularly alkaloids, which are produced by 10–20% of all higher plants [1], are often found in their cationic form in environmental compartments. While organic carbon is recognized to predominately affect any organic compound's phase distribution in soils, clays are one of the major geosorbents in soils and sediments exhibiting potential for sorption of cationic organic chemicals in particular [10, 11]. As demonstrated in a preceding study, sorption of the majority of investigated alkaloids to soil organic carbon (SOC) is dominated by cation exchange processes on negatively charged functional groups [12]. Electrostatic interactions such as cation exchange and sorption to geosorbents other than SOC are rarely integrated in soil sorption models commonly used in environmental exposure assessment. Thus, these models often underestimate sorption of cationic chemicals in soils [13]. However, a model that was specifically developed for cationic organic chemicals suggests that in some cases more than 90% of a cation's overall sorption affinity can be attributed to clay minerals [10, 14]. As an illustrative example, sorption of the indole alkaloid strychnine (STY, base $pK_a = 8.37$) was previously observed to directly correlate with soil clay content rather than with SOC content [15]. In combination, these results highlight the fact that the role of clay minerals in determining the phase distribution of protonated phytotoxins and their mobility in soil may be substantially undervalued.

For SOC, we obtained systematic sorption data for 51 largely protonated N-containing phytotoxins as well as 41 neutral or deprotonated compounds. Thus, we were able to pinpoint structural moieties majorly affecting phytotoxin sorption behavior by using continuous flow column sorption experiments as a high-throughput alternative to traditional batch sorption experiments [12]. The aim of this study was to (1) quantify sorption affinities of a selection of these previously investigated phytotoxins to the two common clay minerals kaolinite and montmorillonite by applying the same methodology as in our previous study; (2) gain further insights into how phytotoxin structural features affect sorbate-sorbent interactions, and (3) compare phytotoxins' sorption affinities to model clay with those to model SOC as a basis to evaluate the relative importance of these major geosorbents in soils.

Experimental section

Sorbates, sorbents and solutions

A set of 52 phytotoxins previously studied with regard to their sorption behavior to SOC [12] was chosen to investigate sorption to clay minerals under environmentally relevant conditions. Selected sorbates were representatives of the major phytotoxin classes alkaloids, terpenoids or steroids and were investigated as predominantly

protonated or neutral species (Table 1, Additional file 1: Table S2). While for the protonated sorbates (i.e., 60% of investigated compounds) cation exchange was identified as the major sorption mechanism to SOC, sorption of the chosen neutral compounds was dominated by either hydrophobic partitioning or complexation reactions.

Stock solutions of phytotoxins (Phytolab, Vestenbergsgreuth, Germany or Sigma-Aldrich, Buchs, Switzerland) as well as solutions of sodium nitrate (NaNO3) and thiourea (CH4N2S, both Sigma-Aldrich) as conservative tracers were prepared in methanol (HPLC grade, Sigma-Aldrich) and stored at - 20 °C. Aqueous eluents used in retention measurements were prepared by dissolving calcium chloride (CaCl2 x H2O, Sigma-Aldrich) in ultrapure water (Milli-Q system, Merck Millipore, Darmstadt, Germany) to get a solution of 5 mM CaCl2, and air-equilibrated for at least 24 h to reach a stable pH of 6 \pm 0.1. Phytotoxin test solutions for sorption experiments (0.01–20 mg $\rm L^{-1})$ were prepared by dilution with the aqueous eluent while ensuring that the methanol concentration remained below 2% in all cases.

Two clay minerals with different cation exchange capacity (CEC) were purchased from The Clay Minerals Society (Chantilly, USA): The non-expandable two-layer clay kaolinite with a CEC of 20 mmol kg⁻¹, a surface area of 10.05 ± 0.02 m² g⁻¹ and 5% permanent charge; and the expandable, three-layer clay montmorillonite with a CEC of 844 mmol kg⁻¹, a surface area of 83.79 ± 0.22 m² g⁻¹, 95% permanent charge and Ca²⁺ as major exchange cation [16]. Both clays were sieved to obtain a homogenous particle size of 32-63 µm (stainless steel sieves ISO 3310-1, RETSCH, Haan, Germany).

Continuous flow column sorption experiments

Sorption affinities of phytotoxins to clay minerals were determined by continuous flow column sorption experiments using manually packed HPLC columns [12]. This approach was shown to be directly comparable to traditional batch sorption experiments on numerous occasions [17, 18]. Empty HPLC columns (stainless steel, length 14 mm, inner diameter 3 mm) and additional equipment (pre-columns, connecting nuts, stainless steel sieves of 3 μm pore size, glass fiber filters of < 1.5 μm pore size and PTFE sealing rings) were purchased from Bischoff Chromatography (distributed by WICOM International AG, Maienfeld, Switzerland). Clay was diluted with silicon carbide (SiC, particle size 10.8–13.3 μm, ESK-SiC, Frechen, Germany) as quasi-inert material to obtain stationary phases with 1% (montmorillonite) or 20% (kaolinite, w/w) clay in respective columns. The column packing method, further modified from literature [19], is described in detail in Schönsee et al. [12] In short, the material mixture was packed into the empty columns and Schönsee et al. Environ Sci Eur (2021) 33:36 Page 3 of 14

Table 1 List of analyzed phytotoxins including their respective charge under experimental conditions

ID	Toxin name	Toxin class	Sum formula	CAS	Charge at pH 6 ^a
AIL	Ailanthone	Terpenoid	C ₂₀ H ₂₄ O ₇	981-15-7	0
ANI	Anisatin	Terpenoid	$C_{15}H_{20}O_{8}$	5230-87-5	0
BAC	Baccatin III	Terpenoid	$C_{31}H_{38}O_{11}$	27548-93-2	0
BIC	(+)-Bicuculline	Alkaloid	$C_{20}H_{17}NO_{6}$	485-49-4	+
3RU	Brucine	Alkaloid	$C_{23}H_{26}N_2O_4$	357-57-3	+
CAF	Caffeine	Alkaloid	$C_8H_{10}N_4O_2$	58-08-2	0
CEV	Cevadine	Alkaloid	$C_{32}H_{49}NO_{9}$	62-59-9	+
CIN	Cinobufagin	Steroid	$C_{26}H_{34}O_{6}$	470-37-1	0
COL	Colchicine	Alkaloid	$C_{22}H_{25}NO_{6}$	64-86-8	0
CON	Convallatoxin	Steroid	C ₂₉ H ₄₂ O ₁₀	508-75-8	0
COS	(+)-Costunolide	Terpenoid	$C_{15}H_{20}O_2$	553-21-9	0
CUE	Cucurbitacin E	Terpenoid	$C_{32}H_{44}O_{8}$	18444-66-1	0
CYT	(—)-Cytisine	Alkaloid	C ₁₁ H ₁₄ N ₂ O	485-35-8	+
DAB	10-Deacetylbaccatin III	Terpenoid	C ₂₉ H ₃₆ O ₁₀	32981-86-5	0
DIG	Digitoxigenin	Steroid	C ₂₃ H ₃₄ O ₄	143-62-4	0
ECH	Echimidine	Alkaloid	$C_{20}H_{31}NO_{7}$	520-68-3	+
ERU	Erucifoline	Alkaloid	C ₁₈ H ₂₃ NO ₆	40158-95-0	+/0
ERUNO	Erucifoline N-oxide	Alkaloid	C ₁₈ H ₂₃ NO ₇	123864-94-8	
GAL	Galanthamine	Alkaloid	C ₁₇ H ₂₁ NO ₃	357-70-0	+
GEL	Gelsemine	Alkaloid	C ₂₀ H ₂₂ N ₂ O ₂	509-15-9	+
GRA	Gramine	Alkaloid	$C_{11}H_{14}N_2$	87-52-5	+
HEL	Heliotrine	Alkaloid	C ₁₆ H ₂₇ NO ₅	303-33-3	+
HOR	Hordenine	Amine	C ₁₀ H ₁₅ NO	539-15-1	+
ISC	(+)-Isocorydine	Alkaloid	C ₂₀ H ₂₃ NO ₄	475-67-2	+/0
JAC	Jacobine	Alkaloid	C ₁₈ H ₂₅ NO ₆	6870-67-3	+/0
LAS	Lasiocarpine	Alkaloid	$C_{21}H_{33}NO_7$	303-34-4	+
LASNO	Lasiocarpine N-oxide	Alkaloid	C ₂₁ H ₃₃ NO ₈	127-30-0	•
LAT	Lathyrol	Terpenoid	C ₂₀ H ₃₀ O ₄	34420-19-4	0
LYC	Lycopsamine	Alkaloid	C ₁₅ H ₂₅ NO ₅	10285-07-1	+
LYO	Lycorine	Alkaloid	C ₁₆ H ₁₇ NO ₄	476-28-8	+/0
MAT	Matrine	Alkaloid	C ₁₅ H ₂₄ N ₂ O	519-02-8	+
MON	Monocrotaline	Alkaloid	C ₁₆ H ₂₃ NO ₆	315-22-0	+/0
NIC	Nicotine	Alkaloid	$C_{10}H_{14}N_2$	65-31-6	+
NUC	(—)-Nuciferine	Alkaloid	C ₁₉ H ₂₁ NO ₂	475-83-2	+
OLE	Oleandrin	Steroid	C ₃₂ H ₄₈ O ₉	465-16-7	0
PAR	Parthenolide	Terpenoid	C ₁₅ H ₂₀ O ₃	20554-84-1	0
PIC	Picrotoxinin	Terpenoid	C ₁₅ H ₁₆ O ₆	17617-45-7	0
PIL	Pilocarpine	Alkaloid	$C_{15} H_{16} O_6$ $C_{11} H_{16} N_2 O_2$	54-71-7	+
PRO	Protopine	Alkaloid	$C_{11}H_{16}N_2O_2$ $C_{20}H_{19}NO_5$	130-86-9	+
PTB	Pterosin B	Terpenoid	C ₁₄ H ₁₈ O ₂	34175-96-7	0
QUI	(—)-Quinine	Alkaloid	$C_{14}H_{18}O_2$ $C_{20}H_{24}N_2O_2$	130-95-0	+
REC	Retronecine	Alkaloid	C ₈ H ₁₃ NO ₂	480-85-3	+
RET	Retrorsine	Alkaloid	C ₁₈ H ₂₅ NO ₆	480-54-6	+/0
SAT	α-Santonin	Terpenoid		481-06-1	
SEK	Senkirkine	Alkaloid	C ₁₅ H ₁₈ O ₃ C ₁₉ H ₂₇ NO ₆	2318-18-5	o +/o
SEN	Senecionine	Alkaloid	C ₁₉ H ₂₅ NO ₅	130-01-8	+/0
SENNO	Senecionine N-oxide	Alkaloid		13268-67-2	⊤ /∪
SOP	Sophocarpine	Alkaloid	C ₁₈ H ₂₅ NO ₆	6483-15-4	ر
			C ₁₅ H ₂₂ N ₂ O		+
STR	Strophanthidin () Strychning	Steroid	C ₂₃ H ₃₂ O ₆	66-28-4 57-24-0	0
STY	(—)-Strychnine	Alkaloid	$C_{21}H_{22}N_2O_2$	57-24-9	+
VIN	(+)-Vincamine	Alkaloid	$C_{21}H_{26}N_2O_3$	1617-90-9	+
YOH	Yohimbine	Alkaloid	$C_{21}H_{26}N_2O_3$	146-48-5	+

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Table 1 (continued)

exposed to a continuously increasing flow of the aqueous eluent in an Agilent HPLC 1260 system (Santa Clara, USA) for liquid compression. Columns packed with SiC only were prepared as reference columns. Preceding any measurement, each column was equilibrated at least overnight by flushing with the eluent at 0.1 mL min $^{-1}$ to ensure that clay minerals were converted to the homoionic form with Ca $^{2+}$ as major exchange cation. Signal intensities were acquired with diode array detection (DAD, Agilent 1290 DAD) at analyte-specific wavelengths (Additional file 1: Table S3). Measurements were generally executed at a flow rate of 0.1 mL min $^{-1}$ and a fixed analysis temperature of 25 \pm 1 °C. For all phytotoxins, 5 μ L of up to eight different dilutions of each individual sorbate were injected.

Using the continuous flow column sorption setup, quasi-equilibrium sorption coefficients ($D_{\rm clay}$) normalized to the amount of clay in the columns ($m_{\rm clay}$) can directly be derived from retention of the sorbates on the clay columns according to Eq. 1.

$$D_{\rm clay} \left[{\rm L~kg^{-1}} \right] = \left(V_{\rm ret, clay} - V_{\rm ret, SiC} \right) \! /_{\! M_{\rm clay}} \tag{1} \label{eq:Dclay}$$

with V_{ret} being the retention volume obtained by multiplying the dead-time corrected retention time of the sorbate with the constant flow rate of 0.1 mL min⁻¹ for the clay and SiC column, respectively. To account for peak asymmetry resulting from pronounced tailing on the clay columns, primarily caused by the occurrence of various retention mechanisms at once (e.g., hydrophobic partitioning and cation exchange), all retention times were obtained based on the statistical first moment approach [20]. Experimentally derived sorption coefficients were further normalized to the CEC of the respective clay mineral ($D_{\text{kao,CEC}}$, $D_{\text{mont,CEC}}$). Phytotoxin D_{oc} data, obtained under the same experimental conditions and also applying the column sorption setup [12], was normalized to the CEC of the organic reference material (2130 mmol kg $^{-1}$) [21] to obtain $D_{\text{oc,CEC}}$ (Additional file 1: Table S1) for direct comparison with clay sorption

[mmol kg⁻¹]). As sorption was expected to mainly occur on the clay surface or limited edge sites, isotherms were fitted with the Freundlich equation (Eq. 2) [23] defined as:

$$c_s = K_F \times c_w^n, \tag{2}$$

where K_F [mmol¹⁻ⁿ L n kg $^{-1}$] is the Freundlich coefficient and n [-] the Freundlich exponent. The latter directly reflects sorption (non-)linearity. A Freundlich exponent of n=1 corresponds to linear sorption, while n<1 indicates saturation of available high affinity sorption sites. Exponents larger than one are typical for cooperative sorption with sorbate–sorbate in addition to sorbate–sorbent interactions. Additional linear fitting was performed for comparison.

Data quality

Quality control measurements were executed as suggested by Bi et al. [24] and method performance was overall comparable with our preceding study on SOC [12]. Clay columns were firstly evaluated by assessing the retention factor k of analytes at different flow rates between 0.025 and 0.15 mL min⁻¹ (Additional file 1: Figure S1). The retention factor k was calculated as follows:

$$k[-] = \frac{(t_R - t_0)}{t_0},\tag{3}$$

with t_R being the retention time of the analyte, and t_0 the dead-time of the system (retention time of the nonretained tracer). Overall, variations in k were negligible suggesting quasi-equilibrium conditions in the columns (Additional file 1: Figure S1). Sorbate recovery was tested by comparison of peak areas for the clay and SiC columns. For 82% of the phytotoxins, recoveries were in the range of 70-110% with a relative standard deviation of 16% and thus, considered sufficient (Additional file 1: Table S3). Any potential interactions of the sorbates with SiC and/or the whole HPLC system were assessed by measurements on the pure SiC column and were generally negligible with t_R close to those of the non-retained tracers for all sorbates. Repeated measurements of the non-retained tracer NaNO3 and thiourea (Additional file 1: Figure S2) showed that column packing was very stable with retention time shifts smaller than 0.05 min. In general, obtained sorption coefficients $D_{\rm clay}$ were only valid if retention on the column was 50% larger than the sum of the retention times of the tracer and analyte on

 $[\]bar{a}$ o = neutral at pH 6; + = protonated at pH 6 (base $pK_a > 6.5$); +/o = partially protonated/partially neutral (5.5 < base $pK_a < 6.5$)

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the SiC only column. Additionally, all data for which retention on the clay column was below 30% of the tracer were discarded. These data quality criteria, adapted from literature [22], resulted in lower limits of log $D_{\rm mont} \sim 0.5$ and log $D_{\rm kao} \sim 0.1$ for the column setup. If sorbates did not pass both these criteria, sorption coefficients could not be quantified and are given as below the respective lower limit. No upper limit was defined as all tested phytotoxins were in the dynamic range of the method. The previously investigated alkaloid nicotine (NIC) showed an average log $D_{\rm mont,CEC}$ of 2.45 ± 0.03 which is in agreement with the literature value of 2.42 ± 0.22 for NIC sorption to clay minerals [11].

Results and discussion

Average values of experimentally derived sorption coefficients for the phytotoxins investigated in this study are given in Table 2 for both kaolinite (log D_{kao}) and montmorillonite (log D_{mont}). For kaolinite, 26 phytotoxins exceeded the quantification threshold of 0.1 log units, while for montmorillonite, 33 had a sorption coefficient > 0.5. Sorption of protonated phytotoxins to montmorillonite was, on a mass basis, on average 1.7 log units stronger than to kaolinite. Neutral sorbates were poorly retained by any sorbent and rarely exceeded the quantification thresholds (log $D_{\rm mont}$ > 0.5, log $D_{\rm kao}$ > 0.1) with retention being limited to the neutral alkaloids, colchicine (COL) and caffeine (CAF), as well as the steroid oleandrin (OLE) on montmorillonite and the steroid cinobufagin (CIN) on kaolinite. These results confirm the notion that clay minerals present geosorbents that are additionally relevant for the sorption of protonated chemicals in the soil environment, while neutral compounds are hardly affected [10]. The reason behind the retention of the abovementioned neutral sorbates onto one of the sorbents is not fully clear. Pronounced sorption of neutral N-containing compounds to soils was described in previous studies and attributed to complexation to surface bound cations [25]. However, total soil distribution coefficients for neutral benzotriazole to different soils correlated with their respective SOC [25] but not clay content [26]. All four neutral phytotoxins mentioned above contain various keto-, ester- or acetylgroups. Those H-donor and -acceptor functional groups may potentially be involved in interactions with the neutral polar domain of the aluminosilicates [10]. Albeit such functional groups are present in most of the other analyzed neutral phytotoxins as well, their sorption may just be less pronounced and below the operational limit of the method. Besides, the phytotoxins COL, CAF and CIN were among those sorbing strongest to SOC by nonspecific interactions [12]. Such non-specific interactions were shown to even contribute to sorption of cationic aromatic amines to montmorillonite [27]. Therefore, it is likely they occur for neutral sorbates as well.

Sorption (non-)linearity

The experimental results enabled the derivation of semi-quantitative sorption isotherms with three to eight concentration levels for 28 and 26 phytotoxins on montmorillonite and kaolinite, respectively. Isotherms for individual sorbates and compound classes on both clays are illustrated in the electronic supporting information (ESI, Additional file 1: Figure S3–S14). Estimated maximum sorbed concentrations lay more than a factor 100 below the CEC of montmorillonite and reached up to 3.3% coverage of cation exchange sites on kaolinite (Additional file 1: Table S4). The Freundlich equation fitted the experimental data very well, with average correlation coefficients (R²) of 0.97 for montmorillonite (Table 3) and 0.99 for kaolinite (Table 4).

Sorption was observed to be largely linear for the majority of the phytotoxins on montmorillonite with an average *n* of 1.09 ± 0.08 and equally good linear fit results (Table 3). Due to the very low maximum surface coverage reached within this study (<0.3%), rather constant sorption affinities were expected for montmorillonite sorption in the investigated concentration range. However, exceptions were observed for the five protonated phytotoxins bicuculline (BIC), gramine (GRA), hordenine (HOR), nuciferine (NUC) and protopine (PRO) with Freundlich exponents in the range of 1.3 to 2.2 (Table 3). The differences in sorption isotherms in comparison with compounds displaying linear sorption on montmorillonite are visualized in Fig. 1 for the isoquinoline alkaloid PRO in contrast to the linearly sorbing pyrrolizidine alkaloid erucifoline (ERU). While $D_{\rm mont}$ showed a relatively constant increase of c_s as a function of c_w for ERU, D_{mont} of PRO increased markedly in the measured concentration range indicating a sigmoidal isotherm. For these isotherms, a plateau is generally expected at higher sorbent loads, which could not be evaluated in this study due to limited sorbate availability or solubility constraints at high concentrations. As c_s differed markedly for the investigated sorbate set (maximum c_s between 7×10^{-5} and 2 mmol kg^{-1}), one reason for why some sorbates showed n > 1 might simply be that the data reflected different parts of the isotherms (different levels of sorbed concentrations). Nonetheless, when comparing the isotherms of PRO and ERU, for which similar c_s were evaluated, sorption non-linearity was rather pronounced for the former (Fig. 1). Thus, we cannot rule out that similar trends would be seen for all sorbates at substantially higher c_s but we certainly observed variations in behavior for different compound classes at comparable c_s .

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 Table 2
 Mean sorption coefficients of phytotoxins to the clay minerals kaolinite and montmorillonite

ID	$\log D_{\rm kao}$ [L kg ⁻¹]	$\log D_{\rm kao,CEC}$ [L mol ⁻¹]	$\log D_{\rm mont} [L kg^{-1}]$	$\log D_{\text{mont,CEC}} [\text{L mol}^{-1}]$
AIL.	< 0.1		< 0.5	
ANI	< 0.1		< 0.5	
AC	< 0.1		< 0.5	
IC	0.89 ± 0.02 (8)	2.59 ± 0.02	2.82 ± 0.19 (4)	2.90 ± 0.19
RU	1.09 ± 0.11 (8)	2.79 ± 0.11	3.05 ± 0.03 (6)	3.12 ± 0.03
AF	< 0.1		2.08 ± 0.01 (4)	2.15 ± 0.01
EV	0.94 ± 0.07 (4)	2.64 ± 0.07	2.01 ± 0.02 (4)	2.08 ± 0.02
IN	0.23 ± 0.15 (4)	1.93 ± 0.15	< 0.5	
OL	< 0.1		2.07 ± 0.02 (4)	2.14 ± 0.02
NC	< 0.1		< 0.5	
OS	< 0.1		< 0.5	
UE	< 0.1		< 0.5	
ΥT	0.25 ± 0.01 (8)	1.95 ± 0.01	2.13 ± 0.01 (4)	2.21 ± 0.01
AB	< 0.1		< 0.5	
IG	< 0.1		< 0.5	
CH	0.16±0.01 (8)	1.86±0.01	2.20 ± 0.18 (5)	2.27 ± 0.18
-11 RU	0.10±0.01 (d) 0.57±0.10 (8)	2.27 ± 0.10	$2.70 \pm 0.06 (10)$	2.77±0.06
RUNO	<0.1	2.27 ± 0.10	< 0.5	2.77 ± 0.00
AL	0.38 ± 0.02 (8)	2.08 ± 0.02	2.98±0.01 (4)	3.05 ± 0.01
EL	<0.1	2.00 ± 0.02	2.39 ± 0.01 (4)	2.47±0.01
RA	0.21 ± 0.11 (4)	1.91 ± 0.11	2.01 ± 0.08 (4)	2.08 ± 0.08
EL		1.91 ± 0.11		
or Or	< 0.1	2.24 ± 0.05	2.17±0.01 (2)	2.24±0.01
C C	0.54±0.05 (4)	2.24±0.05	2.10 ± 0.08 (4)	2.17±0.08
	0.89 ± 0.05 (8)	2.59 ± 0.05	3.00 ± 0.04 (4)	3.07 ± 0.04
AC .	0.97 ± 0.07 (8)	2.67 ± 0.07	< 0.5	251 1 0 10
AS ASNO	0.37 ± 0.03 (8)	2.07 ± 0.03	2.44±0.19 (8)	2.51±0.19
ASNO	< 0.1		1.19±0.01 (2)	1.26 ± 0.01
AT	< 0.1		< 0.5	104 001
C (C	< 0.1		1.77 ± 0.01 (2)	1.84±0.01
O	< 0.1	0.45 0.40	1.58 ± 0.01 (4)	1.65 ± 0.01
AT	0.75 ± 0.12 (8)	2.45 ± 0.12	2.61 ± 0.01 (3)	2.68 ± 0.01
ION	0.26 ± 0.16 (4)	1.96±0.16	2.15 ± 0.23 (5)	2.22 ± 0.23
IC	0.60 ± 0.14 (4)	2.30 ± 0.14	2.38 ± 0.03 (4)	2.45 ± 0.03
UC	1.13 ± 0.01 (8)	2.83 ± 0.01	2.91 ± 0.09 (8)	2.98 ± 0.09
LE	< 0.1		1.37 ± 0.06 (8)	1.45 ± 0.06
AR .	< 0.1		< 0.5	
C	< 0.1		< 0.5	
L	0.74 ± 0.09 (4)	2.43 ± 0.09	2.67 ± 0.03 (3)	2.75 ± 0.03
RO	1.17±0.01 (8)	2.86 ± 0.01	2.96 ± 0.14 (8)	3.04 ± 0.14
ГВ	< 0.1		< 0.5	
UI	0.61 ± 0.11 (8)	2.31 ± 0.11	2.33 ± 0.05 (8)	2.40 ± 0.05
EC	< 0.1		1.63 ± 0.01 (2)	1.71 ± 0.01
T	1.02 ± 0.09 (4)	2.72 ± 0.09	2.40 ± 0.02 (3)	2.48 ± 0.02
AT	< 0.1		< 0.5	
K	0.14 ± 0.09 (4)	1.84 ± 0.09	2.41 ± 0.21 (6)	2.49 ± 0.21
EN .	0.89 ± 0.10 (8)	2.59 ± 0.10	2.81 ± 0.01 (4)	2.88 ± 0.01
ENNO	< 0.1		< 0.5	
OP	0.45 ± 0.07 (4)	2.15 ± 0.07	2.53 ± 0.01 (4)	2.60 ± 0.01
ΓR	< 0.1		< 0.5	
ГΥ	1.11 ± 0.08 (4)	2.81 ± 0.08	2.96 ± 0.03 (4)	3.04 ± 0.03
IN	< 0.1		2.09 ± 0.05 (4)	2.16 ± 0.05
ОН	0.82 ± 0.02 (4)	2.52 ± 0.02	NA	

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Table 2 (continued)

 $\log D_{\rm kao} =$ sorption coefficient to kaolinite; $\log D_{\rm mont} =$ sorption coefficient to montmorillonite; subscript CEC depicts the sorption coefficients normalized to the cation exchange capacity of the different clay minerals; values < 0.1 or < 0.5 are below the operational range of the method; errors are expressed by the standard deviation of experiments; numbers in parentheses = the number of individual retention measurements with quantifiable sorption coefficient for each phytotoxin; for explanation of abbreviations in column "ID" see Table 1; NA = not analyzed

Table 3 Fitting parameters for semi-quantitative isotherms describing sorption of phytotoxins to montmorillonite

ID	Freundlich fit (montmorillonite)			Linear fit (montmorillonite)	
	$\log K_F$	n	R^2	log K _L	R^2
BIC	5.58 ± 0.31 (4)	1.70 ± 0.08	1.00	3.04 ± 0.13 (4)	0.97
BRU	2.84 ± 0.23 (6)	0.95 ± 0.06	0.98	3.03 ± 0.12 (6)	0.99
CEV	1.52 ± 0.22 (4)	0.87 ± 0.06	0.99	1.96 ± 0.01 (4)	0.99
COL	1.85 ± 0.06 (4)	0.93 ± 0.02	1.00	2.04 ± 0.01 (4)	1.00
CYT	2.07 ± 0.02 (4)	0.97 ± 0.01	1.00	2.13 ± 0.01 (4)	1.00
ECH	2.02 ± 0.44 (3)	0.99 ± 0.13	0.98	2.02 ± 0.04 (3)	0.98
ERU	2.66 ± 0.01 (8)	0.98 ± 0.00	1.00	2.72 ± 0.00 (8)	1.00
GAL	3.07 ± 0.03 (4)	1.03 ± 0.01	1.00	3.00 ± 0.20 (4)	1.00
GEL	2.29 ± 0.05 (4)	0.97 ± 0.02	1.00	2.38 ± 0.04 (4)	1.00
GRA	2.84 ± 0.20 (4)	1.33 ± 0.08	0.99	2.14 ± 0.34 (4)	0.97
HOR	5.09 ± 0.26 (4)	2.20 ± 0.10	1.00	2.41 ± 0.61 (4)	0.96
ISC	3.34 ± 0.37 (4)	1.09 ± 0.10	0.98	3.02 ± 0.12 (4)	0.99
LAS	2.55 ± 0.47 (6)	1.07 ± 0.14	0.94	2.30 ± 0.07 (6)	0.94
LYO	1.60 ± 0.00 (4)	1.01 ± 0.00	1.00	1.58 ± 0.01 (4)	1.00
MAT	2.49 ± 0.17 (3)	0.96 ± 0.05	1.00	2.59 ± 0.08 (3)	0.99
MON	1.59 ± 2.82 (3)	0.86 ± 1.02	0.42	1.91 ± 0.74 (3)	0.60
NIC	2.20 ± 0.06 (4)	0.93 ± 0.02	1.00	2.34 ± 0.30 (4)	1.00
NUC	4.07 ± 0.13 (8)	1.33 ± 0.04	1.00	3.01 ± 0.07 (8)	1.00
OLE	0.76 ± 0.16 (8)	0.81 ± 0.05	0.98	1.24 ± 0.01 (8)	0.99
PIL	2.09 ± 0.01 (3)	0.82 ± 0.00	1.00	2.58 ± 0.02 (3)	1.00
PRO	4.35 ± 0.28 (8)	1.33 ± 0.07	0.99	3.12 ± 0.19 (8)	0.90
QUI	2.67 ± 0.14 (8)	1.11 ± 0.04	0.99	2.38 ± 0.11 (8)	0.99
RET	2.64 ± 0.49 (3)	1.07 ± 0.13	0.98	2.45 ± 0.04 (3)	0.98
SEK	2.45 ± 0.01 (4)	0.98 ± 0.00	1.00	2.50 ± 0.00 (4)	1.00
SEN	2.84 ± 0.11 (4)	1.01 ± 0.03	1.00	2.82 ± 0.01 (4)	1.00
SOP	2.43 ± 0.03 (4)	0.97 ± 0.01	1.00	2.51 ± 0.04 (4)	1.00
STY	3.33 ± 0.04 (4)	1.10 ± 0.01	1.00	3.01 ± 0.09 (4)	1.00
VIN	2.62 ± 0.09 (4)	1.15 ± 0.03	1.00	2.16 ± 0.03 (4)	0.99

 K_E [mmol¹⁻ⁿ Lⁿ kg⁻¹] = Freundlich coefficient; n [-] = Freundlich exponent; R^2 = correlation coefficient; K_L [L kg⁻¹] = linear coefficient; numbers in parentheses specify the number of individual dilutions with quantifiable sorption coefficient for each phytotoxin

Overall, investigated concentrations corresponded to an average c_s of 32 mg kg $^{-1}$ (0.1 mmol kg $^{-1}$) and c_w of 2 mg L $^{-1}$ (0.01 mmol L $^{-1}$). From an environmental relevance point of view, both these concentrations far exceeded any field observations for soil or aqueous concentrations of phytotoxins. As an example, the pyrrolizidine alkaloid ERU exhibited concentrations of

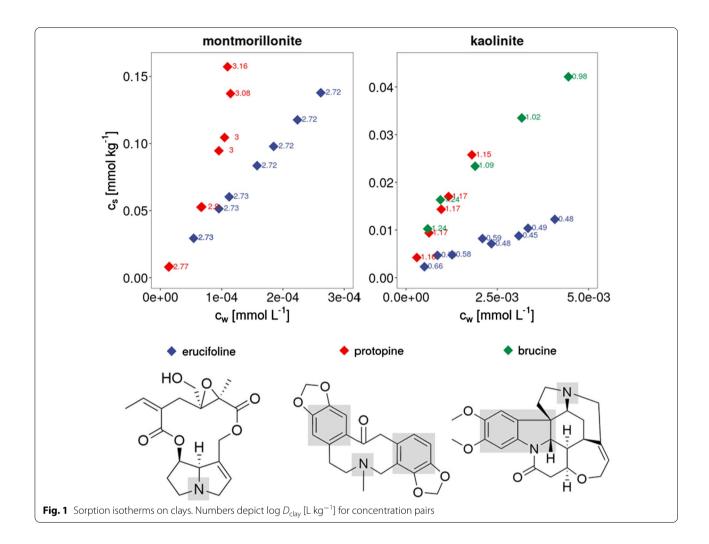
Table 4 Fitting parameters for semi-quantitative isotherms describing sorption of phytotoxins to kaolinite

ID	Freundlich fit (kaolinite)			Linear fit (kaolinite)	
	log K _F	n	R^2	$\log K_L$	R^2
BIC	0.61 ± 0.03 (8)	0.91 ± 0.01	1.00	0.85 ± 0.06 (8)	1.00
BRU	0.36 ± 0.09 (8)	0.72 ± 0.03	0.99	0.93 ± 0.45 (8)	0.98
CEV	0.70 ± 0.07 (4)	0.94 ± 0.02	1.00	0.85 ± 0.46 (4)	0.99
CIN	-0.81 ± 0.06 (4)	0.61 ± 0.02	1.00	0.00 ± 0.04 (4)	1.00
CYT	0.25 ± 0.02 (8)	1.00 ± 0.01	1.00	0.25 ± 0.02 (8)	1.00
ECH	0.06 ± 0.05 (8)	0.96 ± 0.02	1.00	0.13 ± 0.02 (8)	1.00
ERU	-0.21 ± 0.18 (8)	0.72 ± 0.06	0.95	0.40 ± 0.23 (8)	0.95
GAL	0.35 ± 0.01 (8)	0.97 ± 0.01	1.00	0.36 ± 0.01 (8)	1.00
GRA	1.61 ± 0.20 (4)	1.72 ± 0.10	0.99	0.44 ± 0.36 (4)	0.97
HOR	0.88 ± 0.03 (4)	1.19 ± 0.02	1.00	0.61 ± 0.03 (4)	1.00
ISC	0.64 ± 0.06 (8)	0.90 ± 0.02	1.00	0.84 ± 0.08 (8)	1.00
JAC	0.63 ± 0.19 (8)	0.91 ± 0.02	0.98	0.92 ± 0.45 (8)	0.98
LAS	0.09 ± 0.04 (8)	0.89 ± 0.05	1.00	0.31 ± 0.04 (8)	1.00
MAT	-0.43 ± 0.06 (8)	0.57 ± 0.02	0.99	0.47 ± 0.19 (8)	0.98
MON	-0.44 ± 0.43 (4)	0.67 ± 0.21	0.84	0.04 ± 0.23 (4)	0.92
NIC	-0.07 ± 0.01 (4)	0.69 ± 0.01	1.00	0.40 ± 0.17 (4)	0.99
NUC	1.09 ± 0.04 (8)	0.99 ± 0.02	1.00	1.13 ± 0.18 (8)	1.00
PIL	0.10 ± 0.04 (4)	0.75 ± 0.01	1.00	0.60 ± 0.16 (4)	1.00
PRO	1.17 ± 0.01 (8)	1.00 ± 0.01	1.00	1.17 ± 0.06 (8)	1.00
QUI	0.25 ± 0.07 (8)	0.84 ± 0.03	0.99	0.49 ± 0.07 (8)	1.00
RET	0.11 ± 0.02 (4)	0.72 ± 0.01	1.00	0.86 ± 0.29 (4)	1.00
SEK	-0.52 ± 0.03 (4)	0.75 ± 0.01	1.00	-0.01 ± 0.06 (4)	0.99
SEN	-0.03 ± 0.09 (8)	0.71 ± 0.03	0.99	0.71 ± 0.35 (8)	0.97
SOP	-0.03 ± 0.01 (4)	0.79 ± 0.01	1.00	0.33 ± 0.07 (4)	1.00
STY	0.41 ± 0.03 (4)	0.73 ± 0.01	1.00	0.96 ± 0.58 (4)	0.99
YOH	0.80 ± 0.09 (4)	0.99 ± 0.03	1.00	0.80 ± 0.26 (4)	1.00

 K_F [mmol¹⁻ⁿ L^n kg^{-1}] = Freundlich coefficient; n [-] = Freundlich exponent; R^2 = correlation coefficient; K_L [L kg^{-1}] = linear coefficient; numbers in parentheses specify the number of individual dilutions with quantifiable sorption coefficient for each phytotoxin

32 µg kg $^{-1}$ in topsoil and 23 µg L $^{-1}$ in an adjacent pond [28]. Therefore, sorption in an even lower concentration range would be more representative for actual environmental conditions. To evaluate such concentrations, lower method detection limits would be necessary as signal intensities using DAD detection were just too low to properly differentiate peaks from background noise. Thus, presented results mainly highlight differences in sorbate behavior in general.

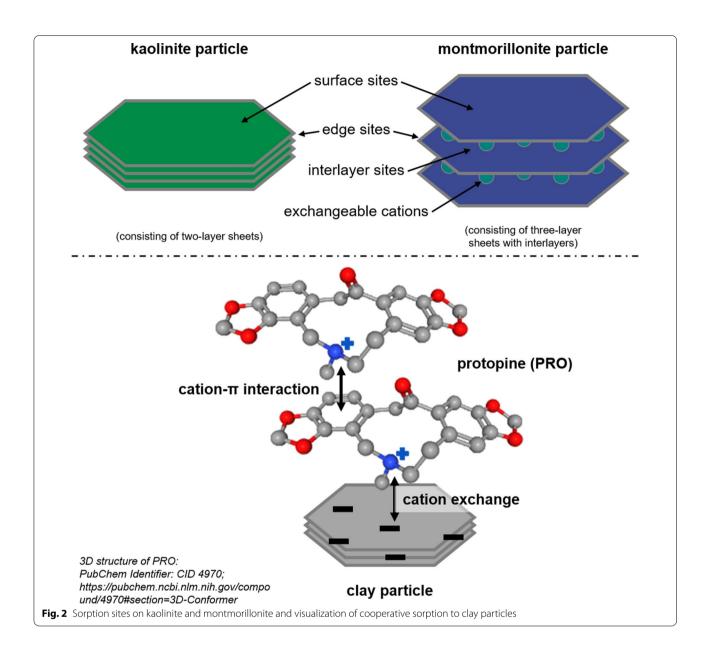
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Freundlich exponents n>1 as well as the isotherm shape shown for PRO (Fig. 1) were indicative of cooperative sorption suggesting enhanced sorbate affinities with increasing surface coverage [11, 29]. Similar ranges for n and sigmoidal isotherms were obtained in previous studies on cationic aromatic amines and clay minerals with interlayers in particular [29, 30]. All compounds that showed this behavior in our study are rather planar, contain tertiary amine functionalities that were protonated under the experimental conditions and additionally possess aromatic moieties in proximity to the charge. It was previously suggested that intermolecular cation- π -, π - π - and van der Waals interactions potentially increase sorption affinities of such cationic aromatic amines to montmorillonite clays (Fig. 2); particularly cation- π -interactions were shown to predominate when just over 0.3% of cation exchange sites were occupied [29]. Based on estimated sorbed concentrations for the compounds with pronounced sorption non-linearity, experimental results presented here are still below this cutoff

(e.g., HOR=0.1%) and also below the proposed limit for sorbate-sorbate interactions of 1 mmol kg⁻¹ (e.g., $HOR = 0.7 \text{ mmol kg}^{-1}$) [11]. However, isotherms indicate that even at these low concentrations, intermolecular cation- π -interactions may occur between the π -systems of a sorbate on the clay surface and the protonated amine group of another sorbate in the vicinity (Fig. 2). To some extent the observed cooperative sorption behavior at lower percentages of cation exchange sites covered could be due to the molecular differences of the sorbates. Phytotoxins investigated in this study are structurally more complex and larger than the comparably simple previously studied aromatic substituted anilines and benzylamines [29]. This tendency may lead to an occupation of available sorption sites at lower concentrations in the case of phytotoxins. An indication for this is given by the fact that N,N-dimethylaniline already reached the range of cooperative sorption when just over 0.3% of cation exchange sites were occupied, while aniline did so only above 2.3% [29].

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In contrast, the average Freundlich exponent for kaolinite was $n\!=\!0.87\!\pm\!0.03$ (Table 4) and the linear fit gave on average larger errors for the sorption coefficients than the Freundlich fit. The maximum surface coverage was reached for the amaryllidaceae alkaloid galanthamine (GAL) with 3.3%, while the average coverage remained low at 0.15% (Additional file 1: Table S4). Whereas ten sorbates showed a Freundlich exponent between 0.9 and 1, indicative of linear sorption (e.g., PRO, ERU, Fig. 1), 14 out of the 26 quantified phytotoxins exhibited a Freundlich exponent below the average of 0.87 for kaolinite (Table 4). One example was the indole alkaloid brucine (BRU) depicted in Fig. 1 with an n of 0.72 \pm 0.03. This

pointed towards complete occupation of surface sites at higher sorbate concentrations for the clay mineral with low CEC in comparison with montmorillonite. However, GRA and HOR showed a tendency towards a sigmoidal isotherm with Freundlich exponents of 1.72 ± 0.10 and 1.19 ± 0.02 for kaolinite as well. The percentage of occupied cation exchange sites was slightly higher than for montmorillonite with, e.g., 0.1-0.5% for HOR. Thus, when exceeding a certain threshold of surface coverage, cooperative sorption could also be observed for the non-expandable clay mineral kaolinite. This has previously only been demonstrated for 4-amino-2-methylquinoline sorption with a Freundlich exponent of 1.52 ± 0.01 [11].

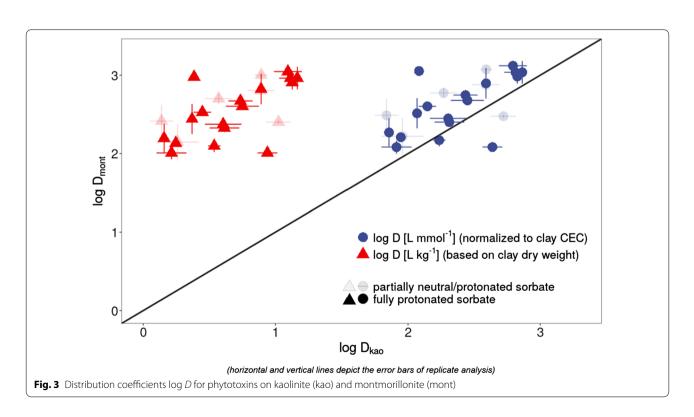
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Normalization of clay sorption affinities to CEC

To rule out CEC as the single driving factor of differences between kaolinite and montmorillonite sorption, obtained log D values were normalized with the respective CEC of each clay mineral. Although it was demonstrated above that sorption was not linear for several sorbates on both clays, normalization was undertaken using the average of all experimentally determined log D values for three to eight individual concentration levels (Table 2). The reason was that linear fitting gave acceptable results with average $R^2 > 0.97$ (Tables 3 and 4) for a majority of the investigated phytotoxins and over the narrow concentration ranges considered in this study. Additionally, taking linearized sorption coefficients for evaluating differences of the two clay minerals allowed direct comparison with trends observed in previous studies on structurally similar compounds [11]. This linearized data was even considered for the development of a sorption model for organic cations predicting their sorption behavior to representative soils with variations of only -0.1 ± 0.4 log units in comparison with directly measured data [14].

As discussed above, sorption affinities of protonated phytotoxins to montmorillonite are substantially stronger than to kaolinite in accordance with the order-of-magnitude higher CEC of the former clay mineral. After normalization to CEC, the variations in $\log D$ between the two clay minerals were minimized to less

than 0.5 log units on average (Table 2, Fig. 3) which is in close agreement with previous studies [11]. Sorption to montmorillonite was still stronger for the majority of the compounds (Table 2, Fig. 3) and must be attributed to structural features of the clay minerals. The mineralogy of different clay minerals was shown to affect sorption behavior of organic cations, particularly for higher order amines or structurally complex sorbates [30]. It was postulated that the position of the negative charge in clay minerals may play a role in sorption of organic cations. The charge in montmorillonite predominantly originates from isomorphic substitutions in the mineral structure when clay lattice metal ions are replaced by other cations of lower charge. As about two-thirds of the negative charge are located in the internal octahedral layers of montmorillonite [31], it exhibits an electrostatic barrier for organic cation uptake [30]. In kaolinite the charge mainly results from pH-dependent (de-)protonation of surface functional groups on edge sites (Fig. 2) [31]. Additionally, sorbates may be more constrained in sorption to interlayer sorption sites on montmorillonite which would affect structurally complex, bulky molecules in particular [30]. The only compound with a sorption affinity to kaolinite that was larger (by more than 0.5 log units) in comparison with montmorillonite after normalization is the steroidal alkaloid cevadine (CEV). With a molecular weight of 591 g mol⁻¹ (Additional file 1: Table S1), it is the heaviest sorbate investigated on both clay minerals.



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This sorbate is almost two times heavier than the average analyte and thus potentially showed a size related restriction to sorb on (otherwise dominating) interlayer sorption sites. It was previously highlighted that the presence of various polar functional groups in addition to the protonated amine functionality mainly affects sorption to clay minerals due to sorbate size rather than polarity, thus limiting interlayer sorption [30]. The steroidal alkaloid CEV with eight polar functional groups (ester-, epoxy-, hydroxyl-groups) represents an excellent example for this observation.

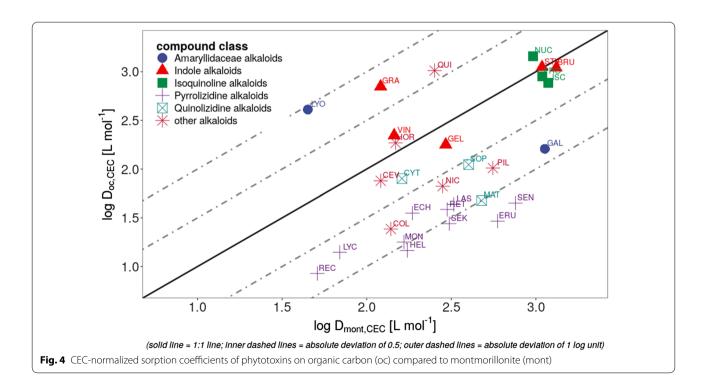
Overall, the extent of sorption for structurally complex molecules with multiple functionalities to clay seems a result of the interplay of charge location and various structural features of the sorbates [30]. The set investigated here on only two clay minerals could not fully shed light on the dominating factors but stresses the need for further investigations in the processes affecting organic cation sorption to geosorbents other than SOC.

Affinities of phytotoxins to clay versus soil organic carbon

To assess the relative affinities of phytotoxins to clay compared to SOC, $D_{\rm oc}$ of 47 compounds [12] determined with the same experimental approach were renormalized for the CEC of the organic reference material used. Out of those, 29 were quantifiable on both montmorillonite and organic carbon. Plotting experimentally derived sorption coefficients for organic carbon (log $D_{\rm oc,CEC}$) against those for montmorillonite (Fig. 4)

shows a relatively clear trend towards enhanced sorption (>0.5 log units) for over 50% (n=16) of analyzed phytotoxins to the clay mineral. These predominantly protonated alkaloids may sorb more strongly to montmorillonite due to the potentially better accessibility of the charge on clay surfaces. Although the CEC of montmorillonite is lower in comparison with SOC [16, 21], the nature and distribution of the charge differs markedly between the two sorbents. Firstly, montmorillonite has a high proportion of permanent charge originating from charge defects in the mineral lattice (95%) in addition to pH-dependent deprotonated hydroxyl-groups on edge sites. In contrast, the charge on SOC results from the deprotonation of various functional groups (e.g., carboxyl-groups) and is fully pH-dependent. Secondly, the charge is therefore more uniformly distributed over the surface of the clay while on SOC it is rather patchy and also affects the structural arrangement of the sorbent [23, 32]. Nevertheless, exceptions to this general preference for montmorillonite were noted.

Nine out of the 29 investigated phytotoxins, i.e., BRU, CEV, gelsemine (GEL), HOR, isocorydine (ISC), NUC, PRO, STY and vincamine (VIN), exhibited similar (i.e., within 0.25 log units) CEC-normalized sorption coefficients to both sorbents. Excluding CEV which showed size constraints potentially hindering interlayer sorption on montmorillonite as discussed above, sorption of the other eight phytotoxins to SOC was attributed to the interplay of cation exchange and further stabilizing



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cation- π - or π - π -interactions [12]. Those stabilizing interactions may be of comparable strength to cooperative sorption demonstrated for HOR, NUC and PRO on montmorillonite. Such a similarity in interaction strength would explain the highly comparable sorption affinities to the two different geosorbents for those phytotoxins with similar structural features (protonated tertiary amines, aromatic moieties). Based on this observation, further structurally similar phytotoxins may show similar trends at higher sorbent loadings that have not been assessed for all sorbates to the same extent.

Only a few phytotoxins, lycorine (LYO), quinine (QUI) and GRA, favored sorption to SOC over that to montmorillonite with sorption coefficients 0.5 to 1 log units larger for the former. The indole alkaloid GRA and the quinolone alkaloid QUI (as well as the above discussed indole alkaloid VIN) are the only N-heteroaromatics in the investigated set. In contrast to the above discussed tertiary amines with aromatic moieties that show cation- π interactions as stabilizing processes, the N-heteroaromatics may rather be stabilized by ternary surface complexation (i.e., including Ca²⁺ as a bridging cation) as suggested for structurally related benzotriazole sorption to SOC [33]. These complexation reactions are also possible for clay minerals, however, the results indicate that if they do occur for phytotoxins on montmorillonite, then at a substantially lower strength than on SOC. Additionally, protonated heteroaromatics are strong π -acceptors that are able to form π - π complexes with aromatic functional groups of SOC [34–36]. Complementary π -moieties cannot be found on bare clay surfaces and thus present another explanation of the substantially weaker sorption affinities of phytotoxin N-heteroaromatics to montmorillonite clay in comparison with SOC. The amaryllidaceae alkaloid LYO showed the in comparison strongest preference for SOC. Based on structural moieties and the observed sorption mechanisms to SOC, similar sorption affinities on clays were expected [12]. In contrast to other structurally similar phytotoxins, LYO has a significantly lower octanol-water partition coefficient (K_{ow}), both based on experiments (log $K_{\text{ow}} = 0.71 \pm 0.05$) [37] and predictions (0.16) [38]. In which way this high polarity affects sorption of LYO to the clay surface but does not hinder relatively strong association with SOC functional groups remains unclear at this point. As a point of interest, the second investigated amaryllidaceae alkaloid GAL showed the exact opposite behavior with markedly enhanced sorption to montmorillonite in comparison with SOC (Fig. 4).

In general, these observations stress once more that a high proportion of organic cation sorption in soils may be assigned to clay minerals rather than organic carbon. Thus, the concept by Droge and Goss, using CEC correction in addition to weight fraction correction of both SOC and clay sorption coefficients in the assessment of organic cation soil sorption [14], seems reasonable. As clay fractions can far exceed the SOC fraction, depending on soil type, such an approach comes closer to real soil conditions than when assigning soil sorption of organic cations to SOC alone. A limitation is that findings cannot be always generalized for major classes of sorbates such as alkaloids for example. Instead individual compound classes have to be investigated in more detail as simple differences in molecular structures (e.g., N-heterocycles like ERU vs. N-heteroaromatics such as GRA) lead to significantly different interactions with geosorbents.

Conclusion

The study of phytotoxin sorption affinities to clay minerals revealed the complexity in sorption of N-containing sorbates that mainly results from their molecular features. The Freundlich fit of obtained isotherms gave good results, but concentration ranges covering several orders of magnitude should be included in future studies to confirm observed tendencies towards cooperative sorption for protonated tertiary amine phytotoxins with aromatic moieties. However, phytotoxins in soil pore water are often found in the ng $\rm L^{-1}$ to maximum low $\rm \mu g \ L^{-1}$ concentration range [39], suggesting that higher concentration ranges are environmentally less relevant.

The exceptions presented additionally highlight that this study only provides first insights into phytotoxin clay sorption behavior and much more systematic data is needed to formulate general trends. For example, X-ray powder diffraction measurements are a powerful technique to further elucidate specific sorption mechanisms [40] and could support future sorption studies. Further, the link between the investigated model sorbents (clay and SOC) and real aged soil constituents still needs to be clearly established, particularly in light of particle coating or aggregation. Nevertheless, the results of our study indicate that clay minerals do not enhance neutral phytotoxin sorption in soils. However, they do indeed play a role in cationic phytotoxin sorption with similar or greater sorption affinities compared to SOC. This role will particularly become relevant for phytotoxins produced by plants growing on clay rich soils. Stronger retention in soils resulting from enhanced sorption to not only SOC but in addition clay minerals would decrease phytotoxin environmental mobility. Conversely, phytotoxins sorbing strongly

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to clay minerals as well, would possess a smaller likelihood to reach water resources but may increase exposure of soil organisms instead.

Abbreviations

SOC: Soil organic carbon; CEC: Cation exchange capacity; STY: Strychnine; SiC: Silicon carbide; $D_{\rm kao}$: Distribution coefficient to kaolinite; $D_{\rm mont}$: Distribution coefficient to montmorillonite; DAD: Diode array detection; NIC: Nicotine; n: Freundlich exponent; $K_{\rm F}$: Freundlich coefficient; $c_{\rm w}$: Aqueous concentration at equilibrium; $c_{\rm s}$: Sorbed concentration at equilibrium; $D_{\rm kao,CEC}$: Cation exchange capacity normalized distribution coefficient to kaolinite; $D_{\rm mont,CEC}$: Cation exchange capacity normalized distribution coefficient to montmorillonite; COL: Colchicine; OLE: Oleandrin; CIN: Cinobufagin; CAF: Caffeine; BIC: Bicuculline; GRA: Gramine; HOR: Hordenine; NUC: Nuciferine; PRO: Protopin; ERU: Erucifoline; BRU: Brucine; GAL: Galanthamine; CEV: Cevadine; GEL: Gelsemine; ISC: Isocorydine; VIN: Vincamine; LYO: Lycorine; QUI: Quinine; $K_{\rm ow}$: Octanol—water partition coefficient.

Supplementary Information

The online version contains supplementary material available at https://doi.org/10.1186/s12302-021-00469-z.

Additional file 1: Table S1. Analyzed phytotoxins and their experimental and predicted properties; Table S2. Molecular structures of all investigated phytotoxins; **Table S3.** Measurement wavelength, recovery and pK_a values of all analyzed phytotoxins; Figure S1. Flow rate effects on retention factor (k) observed for the phytotoxins CAFfeine, CYTisine, GALanthamine, LYcOrine and STRophanthidin on a 1% montmorillonite column; **Figure S2.** Retention time (t_R) comparison of the non-retained tracers nitrate and thiourea (THI) on a 1% montmorillonite column; Table S4. Maximum sorbed concentrations and the percentage of occupied cation exchange sites on both montmorillonite and kaolinite; Figure S3. Amaryllidaceae alkaloid (GALanthamine, LYcOrine) sorption isotherms on kaolinite and montmorillonite; Figure S4. Amine (HORdenine) sorption isotherms on kaolinite and montmorillonite; Figure S5. Cardenolide and bufadienolide (CINobufagin, OLEandrin) sorption isotherms on kaolinite and montmorillonite; Figure S6. Imidazole alkaloid (PILocarpine) sorption isotherms on kaolinite and montmorillonite; Figure S7. Indole alkaloid (BRUcine, GELsemine, GRAmine, STrYchnine, VINcamine, YOHimbine) sorption isotherms on kaolinite and montmorillonite; Figure S8. Isoguinoline alkaloid (BICuculline, ISoCorydine, NUCiferine, PROtopine) sorption isotherms on kaolinite and montmorillonite; Figure S9. Pyrrolizidine alkaloid (ECHimidine, ERUcifoline, JACobine, LASiocarpine, MONocrotaline, RETrorsine, SEnKirkine, SENecionine) sorption isotherms on kaolinite and montmorillonite; Figure S10. Pyridine alkaloid (NICotine) sorption isotherms on kaolinite and montmorillonite; Figure S11. Quinoline alkaloid (QUInine) sorption isotherms on kaolinite and montmorillonite; Figure S12. Quinolizidine alkaloid (CYTisine, MATrine, SOPhocarpine) sorption isotherms on kaolinite and montmorillonite; Figure S13. Steroid alkaloid (CEVadine) sorption isotherms on kaolinite and montmorillonite; Figure S14. Tropolone alkaloid (COLchicine) sorption isotherms on montmorillonite.

Acknowledgements

The authors would like to specifically thank Inge Stockinger (Agroscope), Johanna Keita (Agroscope), Steven T. J. Droge (University of Amsterdam), Florian Metzelder and Inés Rodríguez Leal (Stockholm University) for providing assistance during setup, execution and evaluation of the experiments. ESK-SiC (Frechen, Germany) is acknowledged for donating Silicon Carbide for column experiments. Furthermore, the authors would like to show their appreciation for the continuous support by Kristopher McNeill (ETH Zürich). We also want to thank the four anonymous reviewers whose comments significantly improved our manuscript.

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Authors' contributions

CS performed and evaluated the experiments as well as drafted the manuscript; TB and FW assisted in planning of experiments and revised the text. All authors read and approved the final manuscript.

Funding

This project received funding from the European Union's Horizon 2020 research and innovation program under the Marie Sklodowska-Curie grant agreement no. 722493 (NaToxAq).

Availability of data and materials

The complete dataset supporting the conclusions of this article is included within the article and the Electronic Supporting Information (ESI).

Declarations

Ethics approval and consent to participate

Not applicable.

Consent for publication

Not applicable.

Competing interests

The authors declare that they have no competing interests.

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Received: 30 October 2020 Accepted: 22 February 2021 Published online: 22 March 2021

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