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Mercury distribution in the surface soil of China is potentially driven by precipitation, vegetation cover and organic matter

Zhao-Yang Zhang^{1,3}, Gang Li², Lei Yang¹, Xin-Jun Wang⁴ and Guo-Xin Sun^{1,3*}

Abstract

Background: Understanding the mechanism of Hg accumulation in soil, which is a net Hg sink, at a national scale is important for protecting the environment and improving food safety. The mercury (Hg) distribution in surface soil in China is quite uneven, with relatively high concentrations in southeastern China and low concentrations in northwestern China. The reason for this distribution is inconclusive, especially at the continental scale. In this study, the relative contributions of the key impact factors, including dry and wet deposition, soil organic matter (SOM) and solar radiation to soil Hg, were evaluated.

Results: Wet and dry deposition associated with precipitation and vegetation cover and emissions influenced by SOM are key factors controlling Hg distribution in surface soil. In southeastern China, high levels of wet deposition associated with the South Asia monsoon and dry deposition, enhanced by vegetation canopies, together with low levels of emissions caused by highly vegetated surfaces and solar radiation, are responsible for the high Hg levels in soil (> 0.08 mg/kg). In northeastern China, moderate levels of wet Hg deposition, high levels of dry deposition via throughfall and litterfall, low emissions due to weak solar radiation and high levels of SOM are responsible for high Hg accumulation in soil. In northwestern China, low levels of wet deposition, together with high emissions levels, low vegetation cover (bare soil) and SOM and strong solar radiation, contributed to the low Hg level in the surface soil (< 0.03 mg/kg).

Conclusions: We suggest that wet deposition derived from the Asian monsoon, dry deposition linked to vegetated surfaces and Hg emissions associated with vegetation cover, SOM and solar radiation play key roles in the soil Hg level in China. In other terrestrial environments worldwide, especially in regions with significantly high levels of wet deposition and high amounts of vegetation cover and soil SOM, high Hg concentrations may exist in the surface soil.

Keywords: Soil mercury distribution, Wet and dry deposition, Vegetation cover, Emissions, Soil organic matter, Solar radiation

Background

Mercury (Hg) is an extremely toxic element to humans [1, 2] and ecological receptors [3], with an average crustal abundance of ~ 0.05 mg/kg [4]. Soils and

terrestrial sediments have been documented as acting as net Hg sinks [5] because atmospheric Hg species (e.g., elemental Hg) produced by anthropogenic activities can be subjected to long-distance transport to remote regions and ultimately find their way (wet and dry deposition) into terrestrial or aquatic systems [6]. It has been reported that soil Hg generally accounts for > 90% of the Hg stored in terrestrial ecosystems, and most Hg in soil (> 80%) is from atmospheric inputs [7]. A recent study showed that upland food such as

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Zhang et al. Environ Sci Eur (2020) 32:89 Page 2 of 10

rice is a significant dietary source of Hg in addition to fish consumption [8]. Understanding the mechanism of Hg accumulation in soil at a national scale is important to protect the environment and improve food safety.

It has been reported that the flux of Hg deposition and emissions from soil are controlled by various environmental factors, such as solar radiation, temperature, Hg speciation and content, soil porosity, soil moisture, pH, SOM content, ambient Hg concentration, vegetation cover, and precipitation [9-14]. It is generally accepted that at a small scale (local or regional scale), mining activities or combustion of Hgbearing coal are among the most important factors affecting Hg distribution in soils [15]. However, at continental and subcontinental scales, such as in China, the relevant main drivers of the spatial distribution of soil Hg in the terrestrial environment are poorly understood [16]. It is still unclear at present what factors in addition to lithology play essential roles in the spatial variability in soil Hg distribution at such a large scale and their contribution to soil Hg (Fig. 1).

China is one of the world's largest sources of Hg in the air [18-20]. The Hg distribution in surface soil is mainly from atmospheric Hg deposition [17]. The first soil Hg map in China was obtained from the National Soil Background Survey, which was initiated in the late 1980s and issued in 1994 (Fig. 1a) [21]. Considering the possible influence of rapid industrialization and dramatic urbanization that occurred in China in past decades on Hg emissions and its distribution in surface soil, the soil Hg map (Fig. 1b) was updated by a new geochemical mapping project the National Multi-Purpose Regional Geochemical Survey (NMPRGS) of China, which was initiated in 1999 and completed in 2014 [2, 21]. Comparing the old and new soil Hg maps at the national scale, the spatial distributions of soil Hg were consistent, regardless of the difference in Hg concentration in the same region, although localscale contamination of soil Hg due to gold production or fossil fuel combustion was observed because of unprecedented economic growth in China during the past two decades [22]. Hg distribution in the surface soil in China was uneven before the Chinese economic boom (Fig. 1a), with relatively high concentrations in southeastern and northeastern China and much lower concentrations in northwestern China (Fig. 1). The primary driving factors for Hg distribution in surface soil at a large scale are elusive. The purpose of this study is to elucidate the major factors affecting the distribution of soil Hg and the relative contribution of these factors.

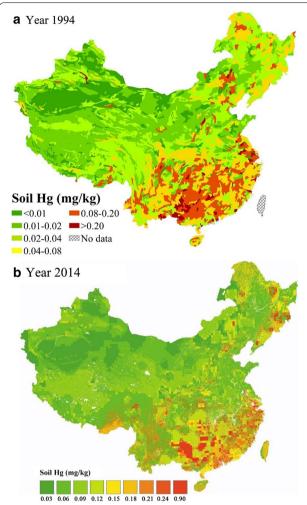


Fig. 1 The spatial distribution of soil Hg in China (mg/kg) surveyed at different periods: **a** map surveyed in the late 1980s from a book titled "The atlas of soil environmental background value in the People's Republic of China" 1994, page 50–51, China Environmental Science Press (modified using ArcGIS Geographic Information Systems software version 10.2, Environmental Systems Research Institute Inc, Redlands, Calif) and **b** updated map surveyed in the 2000s [2]. More details are presented in Ref. [17]. The similar colors in map **a** and **b** refer to distinctive Hg concentration

Materials and methods

The 1994 map of Chinese soil mercury distribution (Fig. 1a) was obtained from the atlas titled "The atlas of soil environmental background value in the People's Republic of China" 1994, page 10–11, China Environmental Science Press, and the 2014 map (Fig. 1b) from a new survey of geochemical mapping project [21] was cited from published literature [2]. The mean annual precipitation dataset was provided by the China Meteorological Data Sharing Service System (Fig. 2a). The Normalized Difference Vegetation Index (NDVI) dataset (Fig. 2b) [23]

Zhang et al. Environ Sci Eur (2020) 32:89 Page 3 of 10

(See figure on next page.)

Fig. 2 The maps of factors and relationships of Hg levels in surface soil with these factors. Spatial distribution of **a** wet deposition in China; **b** NDVI and **c** soil organic matter (created using ArcGIS Geographic Information Systems software version 10.2, Environmental Systems Research Institute Inc, Redlands, Calif). Significant relationships between soil Hg level and **d** precipitation, **f** NDVI and **h** SOM. All data including soil Hg level, precipitation, NDVI and SOM were corrected for variation in sampling intensity by subsampling. Median, range, upper quartile, and lower quartile of soil Hg for each interval of **e** precipitation, **g** NDVI and **i** SOM

was downloaded from MODIS (Moderate Resolution Imaging Spectroradiometer). Data sets have a resolution of 0.05° by 0.05°. NDVI data sets were extracted from per 16-day maximum value composites using a time-series analysis approach [24]. The dataset of SOM (Fig. 2c) were from the National Earth System Science Data Center, National Science and Technology Infrastructure of China (http://www.geodata.cn). The solar radiation distribution dataset (Fig. 4a) was downloaded from SOLARGIS (https://solargis.com/maps-and-gis-data/download/China).

The 1994 map of Chinese soil mercury distribution was digitized with ArcGIS 10.2 (Environmental Systems Research Institute Inc., Redlands, California, USA). We converted the digitized maps of soil mercury distribution, mean annual precipitation, NDVI, SOM, and solar radiation into raster format. The "create random points" tool in ArcGIS 10.2 was used to create sampling points in the maps with a distance of 10 km between every two points. For the Hg values of all obtained points, each point was randomly assigned a value within its particular concentration range to obtain continuous output [24]. All other corresponding values of the obtained points in the different datasets (precipitation, NDVI, SOM, and solar radiation) were extracted by the "extract multi values to points" tool in ArcGIS 10.2. The points with null values were filtered out by the "dplyr" package in the statistical software program "R" (R-Core-Team), and finally, a total of 2827 points were obtained.

Data subsampling to construct sampling bins was utilized to eliminate possible biases caused by uneven data density for different intervals of subsets. The same estimated number of data points was included in each sampling bin. The precipitation dataset was separated into 19 subsets (sampling bins) for every 100 mm/year interval in R, and the last subset contained all points with precipitation amounts higher than 1800 mm/year (Fig. 2d, e). One hundred sample points were randomly selected from each subset of precipitation, and the average value of those points was taken as the representative of this subset (Fig. 2d). The NDVI dataset was divided into 9 subsets based on NDVI values from 0 to 0.9 at intervals of 0.1; 200 sample points were randomly obtained from each sampling bin, and the average value of these 200 sample points represents this subset (Fig. 2f, g). The dataset of SOM was split into 10 subsets at intervals of 1% from 0, and the last subset contained all values > 10%. The average of the 200 randomly sampled points represents this subset (Fig. 2h, i). The dataset of solar radiation was split into 8 subsets. The first subset included the points with solar radiation values less than 1000 kWh/m^2 , and the last subset included all points with values higher than 1800 kWh/m^2 .

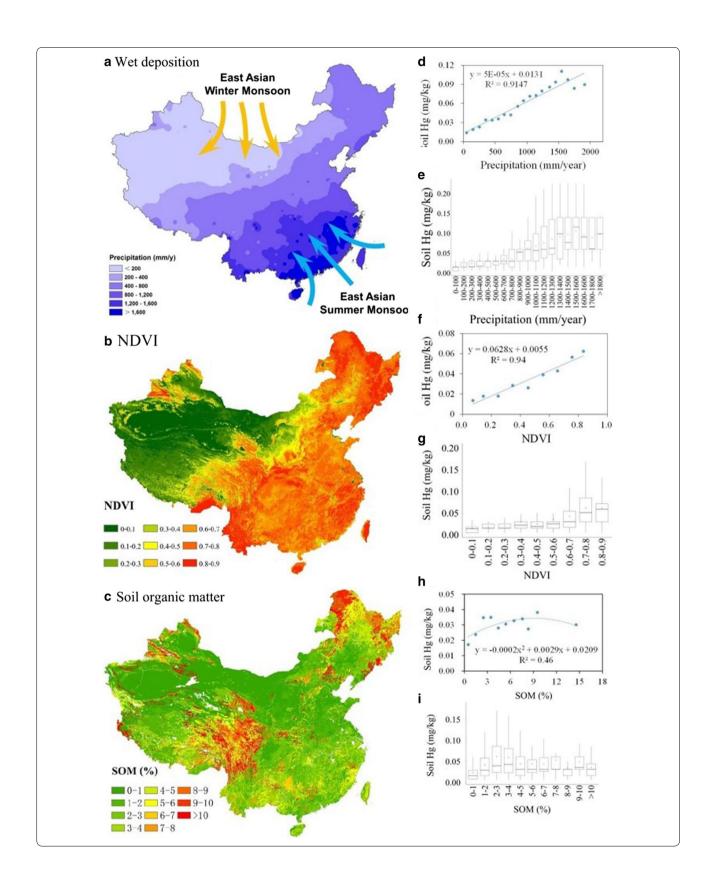
Results and discussion

Mercury input from wet precipitation

A significant linear relationship ($R^2 = 0.91$) between the soil Hg level and precipitation was observed (Fig. 2d). Obviously, a high level of wet deposition was one of the main contributors to high surface soil Hg. The high mean Hg concentration in southeastern China, such as the Yangtze River Delta region, was mainly caused by occasional long-distance transport from domestic source regions and regional anthropogenic emissions due to rapid economic development, which relied heavily on coal combustion for electricity generation [25, 26]. Precipitation derived from the East Asian monsoon is one of the important factors controlling soil Hg distribution in China [27]. In central China, the surface soil Hg concentration was relatively lower than that in southeastern China because the wet deposition in this region was relatively low (400-800 mm). In northwestern China, wet deposition was much less (annual precipitation less than 200 mm), and in some regions, annual precipitation is even less than 100 mm, such as the sand in the Gobi Desert, with no influence from the East Asian summer monsoon. Low Hg concentrations in surface soils (<0.03 mg/kg) were observed (Fig. 1). This trend is also observed in European study results (Additional file 1: Figure S1A). Soil Hg in Europe markedly linearly increased $(R^2 = 0.92)$ accompanied by precipitation (<800 mm/ year). However, higher precipitation (>800 mm/year) did not enhance soil Hg accumulation, and soil Hg accumulation remained constant at high levels (if the last point was omitted due to the geological high Hg background) (Additional file 1: Figure S1A).

A relatively high level of particulate-bound Hg $(0.13-0.18 \text{ ng/m}^3)$ was readily removed through wet deposition [26]. Eastern Chinese soil was largely influenced by anthropogenic Hg emissions from surrounding areas due to its close proximity to the largest anthropogenic Hg emissions. Elevated Hg pollution in the southeastern atmosphere can be expected to have higher deposition

Zhang et al. Environ Sci Eur (2020) 32:89 Page 4 of 10



Zhang et al. Environ Sci Eur (2020) 32:89 Page 5 of 10

Table 1 Summarized Hg wet deposition, dry deposition, emission and net input (deposition minus emission) in soil from different regions, China

| Precipitation (mm/year, i.e. L/m ² /year) | < 400 | 400-800 | | > 800 |
|--|-------------------------------|------------------|------------------|-----------------------------|
| Regions | Northwest | Northeast | Southwest | Southeast |
| Ecosystem | Bare soil or grassland | Forest and shrub | Forest and shrub | Forest, shrub or vegetation |
| Hg wet precipitation (μg/m²/year) | 1.08-4.63 | 8.40 | 0.74–15.9 | 26.1–39.0 |
| Hg dry precipitation (μg/m²/year) | - | 16.5–20.2 | 63.3 | 51.0-66.4 |
| Hg emission (µg/m²/year) | 9.37 | - 1.05 | 9.60 | 51.5-57.8 |
| Net Hg input | − 8.29 ~ − 4.74 | 26.0~29.7 | 54.4~69.6 | 25.6~47.6 |
| References | [47, 59, 60] | [11, 60, 61] | [16, 40, 47, 62] | [59, 61, 63–67] |

in the local soil through precipitation. Most Hg deposition occurred during the monsoon season, with 83% of the oxidized and particulate-bound Hg [28]. Higher Hg concentrations were observed in the sediment samples collected in the Indian estuarine regions during the postmonsoon period than in the pre-monsoon period of the Indian summer monsoon [29], confirming that wet deposition derived from monsoons is one of the major Hg contributors.

The Asian monsoon system controls wet deposition over mainland Asia, including China. During summer monsoons, a strong land—sea pressure gradient is produced by high insolation over continental regions, forcing oceanic winds to converge over southeastern China, bringing oceanic moisture and causing abundant rainfall in this region [30]. The mean annual precipitation decreases gradually from 2000 mm on the southeastern coast to 400–800 mm in central China and then decreases to less than 200 mm in most of northwestern China (Fig. 2a).

The total wet deposition of Hg in the regions of southeastern China was simulated by the global/regional atmospheric heavy metal model (GRAHM) and was consistent with high amounts of precipitation [31]. Monsoons have increased total Hg levels in some regions of the Tibetan Plateau by carrying Hg-enriched air from South Asia [32]. In comparison to nonconvective rainfall events, thunderstorms, which are deep convective rainstorms, are particularly effective at scavenging gaseous oxidized Hg from the boundary layer into the cloud base and result in a nearly 50% increase in Hg deposition [33]. The thunderstorm frequency also greatly contributes to Hg wet deposition. The average annual days of thunderstorms were high (>80 days) in the southern and southwestern regions of China [34], indicating that more Hg wet deposition occurred in these regions, which is consistent with high soil Hg levels (Fig. 1). Excessive Hg input to soil occurred in the monsoon season compared to that in other seasons [28, 35]. Wet deposition typically accounts for 50–90% of the atmospheric Hg entering the environment in some regions [36, 37]. The East Asian summer monsoon generates wind patterns in summer that sweep atmospheric Hg derived from the South China Sea and transport it through the atmosphere and deposit it in terrestrial ecosystems in southeastern China (Fig. 2a).

Mercury input from dry precipitation

In addition to wet deposition, Hg in air also enters terrestrial ecosystems through dry deposition, such as gravity deposition, atmospheric diffusion and other processes [38-40]. Precipitation in central China, such as northeastern China, is moderate (400-800 mm). However, soil Hg levels were relatively high (0.04–0.20 mg/kg, Fig. 1a). This result confirms that dry deposition is also an important contributor to soil Hg accumulation. The significant linear relationship ($R^2 = 0.94$) between the soil Hg level and NDVI at the national scale indicated that vegetation cover has significant effects on Hg accumulation in soil (Fig. 2b, f, g). Typical dry deposition fluxes for elemental Hg over vegetated surfaces were substantially higher than those over non-vegetated surfaces (Fig. 2b, Table 1) [39]. Vegetation is known to contribute to atmospheric Hg as a sink through direct uptake of Hg from the air through stomata on the leaf surface [10, 41]. Upland forest soil acts as a strong sink for Hg, and dry deposition of Hg through litterfall (71.2 µg/m²/year) was the major pathway for Hg loading to the forest catchment in southwestern China [42, 43]. In temperate forested regions such as northeastern China, dry Hg deposition is widely accepted to be approximately two times the amount of wet Hg deposition [44, 45], and uptake of Hg by forest canopies plays a predominant role in soil Hg accumulation [46]. Approximately 72–90% of the total Hg deposition in the forest areas of China was due to dry deposition [27].

Zhang et al. Environ Sci Eur (2020) 32:89 Page 6 of 10

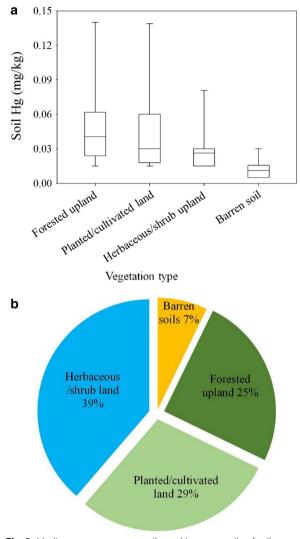


Fig. 3 Median, range, upper quartile, and lower quartile of soil Hg concentrations in different vegetation cover types (**a**). The percentages of Hg in Chinese surface soil with different land cover type (**b**)

In China, soil Hg concentrations are quite different among land cover types and occur in the following order: forested upland (0.055 mg/kg) > planted/cultivated (0.050 mg/kg) > herbaceous upland/shrubland (0.036 mg/kg) > barren soils (0.015 mg/kg), with an average soil Hg concentration that is 3.6 times higher in forests than in barren locations (Fig. 3a). The percentages of Hg contents in surface soil with different land vegetation were different as well (Fig. 3b). Atmospheric Hg deposition fluxes in western China, such as those on the Tibetan Plateau and in Xinjiang Province, ranged from 0.74 to 7.89 μ g/m²/year and were at least one order of magnitude (71.2 μ g/m²/year) lower than that in East Asia (forest cover) [47].

The results of this study showed that different vegetation covers matched the soil Hg distribution in China. In general, Hg concentrations in forest and crop-covered soil (Additional file 1: Figure S2A) were higher than those in grasslands (Additional file 1: Figure S2A). For terrestrial ecosystems, the dry deposition of Hg mainly occurs through the leaves of vegetation absorbing atmospheric Hg from the atmosphere through stomata, and then, Hg enters the forest ecosystem as litterfall [48]. The desert region in northwestern China exhibited the lowest Hg concentrations in soil because of the barren soils. In the western United States, the soil Hg distribution is strongly linked to vegetation greenness with the same order of Hg levels in land cover as that in China, and forested land had more than 2.5 times higher soil Hg levels than that in barren soils [7].

The presence of any soil cover has been found to decrease atmospheric Hg emission fluxes in comparison with bare soils, and even leads to a shift from a net atmospheric emission to a net deposition from the atmosphere [41, 49]. For example, studies were carried out within various types of forest systems across a vast region from the south (South Carolina) to the north (Maine) in United States during the same season. Mercury emission associated with the litter covered forest floor was extremely low [49], fluctuating around 0 ng/ m²/h, with many deposition rather than emission occurring at both daytime and nighttime, irrespective of the forest type, soil type, and variations in weather conditions [10, 49]. Canopy covers in forests and grasslands/ shrublands are partially expected to reduce Hg emissions by limiting soil warming and solar loads via shading the forest floor [31, 50].

Mercury sequestration by soil organic matter

SOM contributes to the accumulation of Hg in soil. Soil Hg concentrations are positively related to soil organic carbon [51, 52], and Hg emissions is inversely correlated with organic content in soils because organic carbon has a high affinity to Hg in soil through functional groups, particularly thiols [7, 53]. In general, with an increase in SOM in soil, the Hg concentration in soil substantially increased (Fig. 2c, h, i). A similar relationship between SOM and Hg concentration was observed in Europe as well (Additional file 1: Figure S1B). In aqueous systems, Hg(II) is reduced to volatile elemental Hg by organic acids, and these reactions appear to be strongly induced by sunlight [54]. These processes do not occur in soils because sunlight cannot easily radiate soil, especially with vegetation cover. Binding of Hg with organic matter inhibits the reduction of Hg(II) to elemental Hg and subsequent emissions from soil. A strong correlation between soil Hg concentrations and organic carbon

Zhang et al. Environ Sci Eur (2020) 32:89 Page 7 of 10

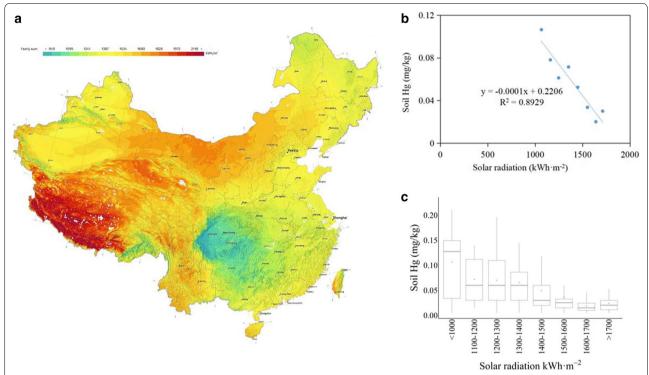


Fig. 4 The maps of solar radiation in China and relationship of Hg levels in surface soil with solar radiation. Spatial distribution of **a** solar radiation in China; significant relationships between soil Hg level and solar radiation (**b**); median, range, upper quartile, and lower quartile of soil Hg for each interval of solar radiation (**c**)

has been observed in the western United States [7]. In northeastern and southwestern China, the distribution of Hg in these regions was geographically coincident with the contents of SOM, especially in northeastern China (Figs. 1 and 2h) [55]. High Hg levels were observed in the northeastern regions, although wet deposition was relatively low (400–800 mm). Soils covered by forests and enrichment in organic matter enhanced Hg deposition by litterfall and throughfall, simultaneously depleting Hg emissions. In highly organic regions (peat bog ecosystems), it is estimated that dry deposition of gaseous elemental mercury is dominant (~80%) over wet deposition (~20%) [56].

Mercury emissions from soils

The solar radiation intensity in China declined from the northwest to the southeast (Fig. 4a), and a strong negative correlation ($R^2 = 0.89$) was observed between the soil Hg level and solar radiation (Fig. 4b), which indicated that increasing solar radiation would accelerate the emissions of Hg from the soil to the atmosphere. It is commonly accepted that gaseous elemental Hg emissions from soil are one of the most important processes of Hg losses in terrestrial environments [16, 57]. Solar radiation, which induces photoreduction, is the most common

and pronounced factor for gaseous elemental Hg emissions. Many studies have identified a positive relationship that is attributed to a photochemically mediated reduction that coverts soil Hg(II) to volatile element Hg [50], although other biotic (microorganisms) and abiotic (organic matter) factors reduced Hg as well [58]. In comparison to direct radiation on bare soil, indirect radiation due to vegetation cover could decrease the role of solar radiation in Hg emissions [31]. Desertified lands (sand desert and Gobi Desert) in northwestern China have bare soils. Direct solar radiation due to poor vegetation cover in northwestern China was one of the contributors to low soil Hg levels.

In addition, the positive association between soil clay content and Hg concentration [51] suggested that atmospheric Hg was hardly sequestrated into deserted and sandy soils with low clay content in northwest China, which has arid and semi-arid climate [55].

Factors controlling soil Hg distribution in China

We hypothesize that the spatial variability in soil Hg levels in China is potentially caused by the balance between wet depositions associated with the East Asian summer monsoon, dry deposition associated with forest/

Zhang et al. Environ Sci Eur (2020) 32:89 Page 8 of 10

vegetation cover and inhibited emissions associated with SOM and solar radiation.

The model of the distribution of soil Hg was established by multiple linear regression as follows:

$$\log(\mathrm{Hg}) = -2.809 + 9.398 \times 10^{-4} \times (P)$$

$$+ 1.1774 \times 10^{-1} \times (\mathrm{NDVI})$$

$$+ 8.879 \times 10^{-2} \times \log(\mathrm{SOM})$$

$$- 9.897 \times 10^{-5} \times (\mathrm{SR}),$$

$$P < 0.01, R^2 = 0.51,$$

where Hg is the soil Hg level (mg/kg) in the 0–20 cm surface soil, P is the average annual precipitation (mm), SOM is the SOM (%), NDVI is the Normalized Difference Vegetation Index, and SR is solar radiation (kWh/m²). Both simple regression analysis of soil Hg with the first three variables and multiple linear regression analysis showed positive relationships, and solar radiation showed a negative relationship between the soil Hg level and those variables. Statistical analysis showed that the relative contributions of the four variables were ranked as follows: precipitation (49.4%) > NDVI (23.4%) > solar radiation (17.9%) > SOM (9.3%).

Conclusions

We suggest that wet deposition associated with the summer Asian monsoon, dry deposition associated with vegetation cover and inhibited emissions linked with SOM and photoreduction due to solar radiation are the key factors controlling Hg redistribution in terrestrial environments in China. Vegetation cover, SOM and solar radiation are key factors controlling Hg air/soil exchange, such as Hg emissions. Statistical analysis showed that of the factors affecting the soil surface Hg concentration, precipitation was the most important factor (49.4%), followed by vegetation cover (23.4%). The contributions of monsoons and vegetation cover to surface soil Hg would be helpful to understand the potential ecosystem and human health risks from Hg and to predict the trend in soil Hg in the future.

Supplementary information

Supplementary information accompanies this paper at https://doi.org/10.1186/s12302-020-00370-1.

Additional file 1: Figure S1. The relationships of soil Hg concentrations in Europe with (A) European precipitations and (B) soil organic carbon (SOC). The soil Hg data, precipitation and soil organic carbon used in this study are from the Geochemical Mapping of Agricultural Soils (GEMAS) project, which was carried out by the EuroGeoSurveys (EGS) Geochemistry Expert Group. The precipitation dataset Europe was separated into 19 subsets (sampling bin) with every 100 mm/year interval in R, and the last subset contains all points with precipitation greater than 2200 mm/year. One hundred sample points were randomly selected from each subset

and the average value of those points was taken as the representative of this subset. The soil organic carbon (SOC) dataset was divided into 13 subsets at interval of 1 g/kg from 0 to 10 g/kg, 10–13 g/kg, 13–26 g/kg and last subset contains all points with SOC more than 26 g/kg. The average of the 200 randomly sampling points obtained for each sampling bin represents this subset. **Figure S2.** (A) vegetation cover (created using ArcGIS Geographic Information Systems software version 10.2, Environmental Systems Research Institute Inc, Redlands, Calif); (B) soil Hg map surveyed in 1994.

Abbreviations

Hg: Mercury; SOM: Soil organic matter; NDVI: Normalized Difference Vegetation Index; NMPRGS: National Multi-Purpose Regional Geochemical Survey; P: Average annual precipitation; SR: Solar radiation.

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Authors' contributions

ZYZ, GXS designed research and wrote the paper; GL, LY contributed data analysis; and XJW has substantively revised the manuscript. All authors read and approved the final manuscript.

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Availability of data and materials

The datasets used and analyzed during the current study are available from the corresponding author on reasonable request.

Ethics approval and consent to participate

Not applicable.

Consent for publication

Not applicable.

Competing interests

The authors declare no competing financial interest.

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Zhang et al. Environ Sci Eur (2020) 32:89 Page 10 of 10

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